ASME231	Atmospheric Thermodynamics	NC A&T State U
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Lecture 15 The Equation of State for Moist Air

(Sec.4.3 of Hess – Equation of State for Moist Air)

A moist air can be treated as a mixture of dry air and water vapor, i.e.

Moist air = dry air + water vapor

We have already learned about how to compute the mean molecular weight of dry air (M_d) from Lecture 2,

$$\overline{M} = \frac{m_1 + m_2 + m_3 + \dots}{\frac{m_1}{M_1} + \frac{m_2}{M_2} + \frac{m_3}{M_3} + \dots} = \frac{1}{(m_1/m)/M_1 + (m_2/m)/M_2 + (m_3/m)/M_3 + \dots}$$

or
$$\overline{M} = 1/\sum \left(\frac{m_i/m}{M_i}\right)$$
, (2.16)

$$M_d = 1/[(m_{\rm N2}/m)/M_{\rm N2} + (m_{\rm O2}/m)/M_{\rm O2} + (m_{\rm Ar}/m)/M_{\rm Ar}]$$

= 1/[0.7551/28+0.2314/32+0.0135/40] = 28.97 kg kmol⁻¹

Thus, the mean molecular weight of the moist air (\overline{M}) can be computed as follows:

The molecular weights of M_d (dry air) and M_v (water vapor)

$$M_d = 28.97 \ kg \ kmol^{-1}, M_v = 18 \ kg \ kmol^{-1},$$

and the respective masses of dry air and water vapor of a moist air, m_d and m_v are used to compute the mean molecular weight of moist air

$$\overline{M} = \frac{m_d + m_v}{m_d / M_d + m_v / M_v}.$$

or in terms of the mass ratio m_v/m_d , which is usually referred to as the mixing ratio ($w=m_v/m_d$):

$$\overline{M} = \frac{1+w}{1/M_d + w/M_v} \,.$$

If we denote M_{ν}/M_d as ε (=18/28.97=0.622), then \overline{M} can be rewritten as

$$\overline{M} = \frac{(1+w)M_d}{1+w/\varepsilon} \,.$$

We can further compute the gas constant for the moist air,

$$\overline{R} = \frac{R^*}{\overline{M}} = \frac{(1 + w/\varepsilon)(R^*/M_d)}{1 + w} = \frac{(1 + w/\varepsilon)R_d}{1 + w}$$

Thus, the equation of state for a moist air may be written,

$$p\alpha = \overline{R}T = R_d \left(\frac{1 + w/\varepsilon}{1 + w}\right)T_{\perp}$$

If we define a new temperature variable,

$$T_{\nu} = \left(\frac{1 + w/\varepsilon}{1 + w}\right)T, \qquad (5.3.1)$$

which is called <u>virtual temperature</u>, then the equation of state for a moist air may be expressed in a similar way as the equation of state for a dry air

$$p\alpha = R_d T_{v_\perp} \tag{5.3.2}$$

The physical meaning of the virtual temperature is the temperature that the dry air would have if its pressure and specific volume were equal to those of a given sample of moist air.

Since $\varepsilon = M_v / M_d = 18/28.97 = 0.622$, and practically $w \ll 1$, Eq. (5.3.1) may be approximated as

$$T_{v} = \left(\frac{1+w/\varepsilon}{1+w}\right)T \approx (1+w/\varepsilon)(1-w)T \approx (1+(1/\varepsilon-1)w-w^{2}/\varepsilon)T \approx (1+0.61w)T$$
$$T_{v} = (1+0.61w)T \qquad (5.3.3)$$

Next, we need to express *w* in terms of state variables and known constants, so that it <u>can be measured</u> in the atmosphere.

By definition,

$$w = m_v / m_d = \rho_v / \rho_d \tag{5.3.4}$$

Using the pressures from the equation of state,

$$p_d = \rho_d (R^*/M_d)T,$$
$$e = p_v = \rho_v (R^*/M_v)T,$$

to replace the densities, we obtain,

$$w = \left(\frac{M_v}{M_d}\right) \left(\frac{e}{p_d}\right),$$

or

$$w = \frac{M_{\nu}}{M_{d}} \left(\frac{e}{p-e}\right) \approx \frac{M_{\nu}}{M_{d}} \left(\frac{e}{p}\right) = \varepsilon \left(\frac{e}{p}\right) = 0.622 \frac{e}{p}.$$
 (5.3.5)

Since *w* is usually less than 0.04, as a result, $T_v - T < 7$ K.

Since the equation of state of dry air is

$$p\alpha = R_d T,$$

comparing with (5.3.2), we may conclude that dry air is denser than moist air. This also explains why the moist air rises in the atmosphere.